



28 **Abstract**

29 The Pan-Third Pole (PTP) owns complex geography and demographic features where
30 aerosol roles and their impact cannot be neglected as it jeopardizes both the environment and
31 human health. Therefore, we analyzed spatio-temporal aerosol concentration, the influence of
32 meteorological conditions, and underlying aerosol transport mechanisms over the PTP by
33 leveraging observation, satellite dataset, and model outputs. The observation and model simulation
34 result showed that aerosol concentrations exceeded the world health organization (WHO) and
35 China guideline values in most of the locations. This study revealed distinctive seasonality with
36 the highest and lowest aerosol concentrations during the winter and summer seasons, respectively,
37 which could be favored by meteorological conditions and emissions from biomass burning. In
38 response to higher aerosol concentrations, the maximum aerosol optical depth (AOD) values were
39 observed over the major hotspot regions however, interestingly summer high (AOD > 0.8) was
40 observed over the Indo Gangetic Plain (IGP) in South Asia. The columnar aerosol profile indicated
41 that the higher aerosol concentrations were limited within 1-2 km elevation over the densely
42 populated regions over South Asia and Eastern China. However, the significant aerosols
43 concentrations found to be extended as high as 10 km could potentially be driven by the deep
44 convection process and summer monsoon activities. Regionally, the integrated aerosol transport
45 (IAT) for black carbon (BC) and organic carbon (OC) was found to be maximum over SA.
46 Noticeable OC IAT anomaly (~5 times > annual mean) found during spring that was linked with
47 the biomass burning events. Yet, the dust transportation was found to be originated from the arid
48 land and deserts that prolonged especially during summer followed by spring seasons. This study
49 highlights the driver mechanism in aerosol seasonality, transport mechanism, and further motivates
50 the additional assessment into potential dynamic relation between aerosol species, aerosol
51 atmospheric river, and its societal impact.

52

53 **Keywords:** *Aerosols, AOD, IAT, Pan-Third Pole, transport dynamics, WRF-Chem*

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56



57 1. Introduction

58 On a regional to the global scale, the atmospheric aerosol has been a major research topic over
59 the recent decades as aerosol impacts ecosystems, climate, and human health. Either way, direct
60 or indirect interaction aerosols perturb the amount of solar radiation reaching the Earth's surface
61 (Twomey, 1977; Haywood and Boucher, 2000; Charlson et al., 1992). The nexus between poor air
62 quality and adverse health outcomes have been documented by epidemiological studies (Shiraiwa
63 et al., 2017; Ravishankara et al., 2020; Lelieveld et al., 2015; Jerrett, 2015; Hong et al., 2019).
64 Globally, ambient air pollution is a major health risk as several adverse health outcomes are linked
65 with poor air quality (Chen et al., 2018; Manisalidis et al., 2020). Due to exposure to ambient air
66 pollution, 4.2 million premature deaths have been reported worldwide in a year in 2016 (Who,
67 2021). Earlier studies articulate the significant aerosol loadings over Asia and its impact
68 (Ramanathan et al., 2007; Carmichael et al., 2009). The regions with population density in South
69 Asia (SA) and East China (EC) face episodic acute air pollution events that are primarily brought
70 forth by a dramatic increase in emissions from rapid urbanization, industrial expansion, power
71 generation, and transportation sectors (Kumar et al., 2020; Yang et al., 2020). Recent events like
72 a super sandstorm in China (Yin et al., 2021; Liu et al., 2021), haze pollution in Delhi (Jena et al.,
73 2021; Dhaka et al., 2020), and extreme air pollution in Kathmandu (Islam et al., 2020; Putero et
74 al., 2015) again raised the prominent concern among all environmental issues over the region.

75 The Pan-Third Pole (PTP) comprises relatively pristine Himalayas and Tibetan Plateau
76 (HTP) but is it surrounded by arid regions to the west and highly polluted regions like Indo-
77 Gangetic Plain (IGP) to the south. The significant aerosol concentrations found over the
78 background region (i.e., HTP) are mainly driven by long-range transport (Yang et al., 2018; Zhang
79 et al., 2015a; Han et al., 2020; Gabrielli et al., 2020). Aerosols like black carbon (BC) and dust
80 could enhance glacier melting thereby posing threat to the hydrological cycle and water availability
81 (Kang et al., 2010; Menon et al., 2002; Ming et al., 2008; Sarangi et al., 2020; Kang et al., 2019).
82 Thus, the fate of cryospheric bodies also depends on the transport of emissions from upwind source
83 regions. Other than this, the complex interplay of aerosols is associated with hydrological
84 processes. It is reported that an increase in aerosol loading and their effects on radiation and cloud
85 microphysics can lead to a weaker Asian hydrological cycle and reduce monsoon precipitation
86 (Ramanathan et al., 2001; Takahashi et al., 2018). A comprehensive review gives more insights
87 into interactions between Asian aerosols and monsoon (Li et al., 2016b).



88 In light of the aforementioned evidence of aerosol impacts, attempts have been made from
89 the quantification of aerosol physical, chemical, optical, and radiative properties to understand the
90 relation between cause and effect of severe air quality over Asia, of which, the majority of studies
91 are mainly focused on the highly polluted regions in South Asia (Ojha et al., 2020; Moorthy et al.,
92 2005; Kumar et al., 2015; Ghude et al., 2016; Ramachandran et al., 2020b) and East China (Huang
93 et al., 2020; Uno et al., 2020; Dang and Liao, 2019; Ding et al., 2016). However, in a relatively
94 cleaner region (i.e., HTP) such potential impacts have gained attention only recently. To quantify
95 the source-receptor transport mechanism over the HTP, the emission perturbation (Yang et al.,
96 2018; Han et al., 2020), backward-trajectory (Lüthi et al., 2015; Lu et al., 2012), tracer-tagging
97 (Kumar et al., 2015; Zhang et al., 2015b), and adjoint (Kopacz et al., 2011) method has been
98 deployed.

99 Also, more studies have been carried out taking into account the aerosol impact on
100 meteorology and vice versa (Zheng et al., 2015; Vinoj and Pandey, 2022; Yuan et al., 2020; Zhao
101 et al., 2017; Huang et al., 2018; Ojha et al., 2020; Lv et al., 2020). As stated by Kumar et al. (2018),
102 Representative Concentration Pathways (RCP) scenarios project a global improvement in air
103 quality, while it will continue to deteriorate over SA over at least the next two decades. It was
104 found that PM_{2.5} (particulate matter with aerodynamic diameter $\leq 2.5 \mu\text{m}$) will increase across
105 South Asia due to the combined effects of the increase in emissions of air pollutants and
106 meteorological conditions under RCP8.5. Kulkarni et al. (2015) also stated that worsening air
107 quality in coming days over SA continues to breach the WHO guideline values. The same study
108 projected that PM_{2.5} decreases over EC and increases significantly in South Asia and Central Asia
109 under the reference scenario (2030). The national air quality observation database and modeling
110 studies indicate that there is a significant decrease of ambient PM_{2.5} concentrations over China
111 (Ding et al., 2019; Uno et al., 2020; Kanaya et al., 2020; Fan et al., 2020; Kong et al., 2021; Zhang
112 et al., 2019a) which also supported by RCP scenario projection study as well (Li et al., 2016a).
113 Whereas, multiple inter-comparison of Aerosol Optical Depth (AOD) data from 2003 to 2017 also
114 indicated an increasing trend over South Asia while decreasing in East China (Gui et al., 2021; Li,
115 2020). In addition, the PTP region includes the number of deserts and arid regions that are the
116 primary source of natural dust emissions and a significant contributor to ambient particulate matter
117 (Chinnam et al., 2006; Ginoux et al., 2001; Chen et al., 2017). Over the PTP region, numerous
118 systematic efforts have been made for the analysis of dust sources, transport, and impact of dust



119 on radiation budget and glaciers in the region (Kumar et al., 2014; Zhang et al., 2020b; Zhao et al.,
120 2010; Li and Sokolik, 2018; Chen et al., 2017; Nabavi et al., 2017; Parajuli et al., 2019)

121 The PTP is highly vulnerable from a climate and human dimension perspective. However,
122 given its fragile ecosystems of global importance the region has not received the attention it
123 deserves in terms of deteriorating air pollution in and around the PTP. Owing to complex
124 geography from low land, arid, semi-arid to high mountain regions where the aerosol burden from
125 surface to the upper troposphere and lower stratosphere over the PTP region from emissions
126 associated with human and natural activities that is processed under the dynamic and vigorous
127 transport mechanisms leads to intricate aerosol-geography-radiation-cloud-climate interactions
128 over the region that are yet to be understood properly. The majority of past studies focused on East
129 China and South Asia regions. However, synoptic-scale studies covering the PTP are scarce which
130 reflects the necessity of further investigation of aerosol distribution, aerosol climatology, long-
131 range transport, and topographic effect on a seasonal basis across the PTP and quantifying
132 contributions of various mechanisms and processes. Therefore, here we used a combination of
133 observations, state-of-art Weather Research and Forecasting coupled with the Chemistry (WRF-
134 Chem) model, the Hybrid Single-Particle Lagrangian Integrated Trajectory (HYSPLIT) model,
135 and the Image-Processing-based Atmospheric River Tracking (IPART) module. To our
136 knowledge, this paper attempts to offer for the first time the quantification of integrated aerosol
137 transport (IAT) using the IPART module for the PTP region. The main objectives of this study
138 are: (1) to evaluate the performance of the model (WRF-Chem), and (2) to understand the spatio-
139 temporal distribution, vertical profile, meteorological influence, and transport mechanism of
140 aerosols ($PM_{2.5}$, PM_{10} , organic compound (OC), BC, and dust). We anticipate that our study helps
141 to bridge the gap in the understanding of a complex interplay between aerosols-climate and
142 associated impacts over the region of global significance. The paper is organized as follows. The
143 details on model setup and observation data used in the study are described in Sec. 2. The spatio-
144 temporal distributions and transport processes are analyzed in Sec. 3. The major findings from the
145 study are summarized in Sec. 4.

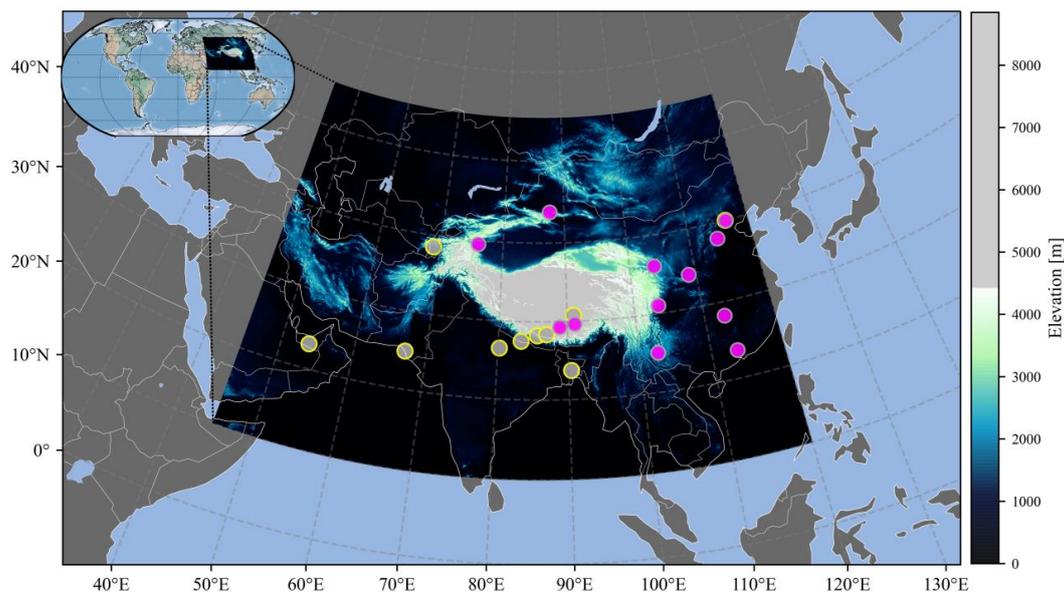
146 **2. Methods**

147 **2.1. WRF-Chem model**

148 In this study, WRF-Chem (version 3.6), a mesoscale three-dimensional Eulerian chemical
149 transport model which enables the feedback between meteorology and chemical processes was



150 used (Skamarock et al., 2005; Grell et al., 1994). We conducted a numerical simulation in a domain
151 with 150×208 horizontal grids with a horizontal resolution of 30 km for the simulation of $PM_{2.5}$,
152 PM_{10} , OC, dust, and meteorological parameters. The model has 40 vertical sigma levels from the
153 surface to the top of the level (50 hPa). The initial and lateral boundary conditions of
154 meteorological forcing were provided by 6 h the National Center for Environmental Prediction
155 final analysis (NCEP/FNL) data on 1×1 grids (<https://psl.noaa.gov/data/>). The chemical initial and
156 boundary conditions were prepared using the mozbc utility tool based on the output of Model for
157 Ozone and Related Chemical Tracers (MOZART) results acquired from the National Center for
158 Atmospheric Research (NCAR) (<https://www.acom.ucar.edu/wrf-chem/mozart.shtml>). We
159 adopted the Morrison-2-moment (Gustafson Jr et al., 2007), the Rapid Radiative Transfer Model
160 (RRTMG) (Iacono et al., 2008; Mlawer et al., 1997), the revised MM5 (Paulson, 1970), the unified
161 Noah (Chen and Dudhia, 2001), the GOCART (Ginoux et al., 2001), and the Yonsei University
162 (YSU) schemes in this study. The simulation runs from January to December 2017, with the first
163 week of January as a spin-up.



164
165 Figure 1. The research domain (dark blue shade) covering the Pan-Tibetan Plateau (PTP) region,
166 with elevation shown in color code. The magenta and gray circular dots illustrate the location of
167 PM ($PM_{2.5}$ and PM_{10}) and the AERONET stations, respectively for which data is used for the
168 model validation.



169 2.2. IPART module

170 IPART is a python module embedded with a new detection algorithm for tracking
171 atmospheric rivers (AR) (Xu et al., 2020). However, in the present study, we utilized this module
172 for extending the AR concept to aerosols considering the zonal and meridian long-range impact
173 on air quality and extremes. To quantify IAT, we leverage the Modern-Era Retrospective analysis
174 for Research and Applications (MERRA-2) assimilated product (Randles et al., 2017) with IPART.
175 Such 2D single-level aerosols diagnostic dataset comes with $0.5^\circ \times 0.625^\circ$ horizontal resolution. The
176 data field is time-stamped with the central time of an hour starting from 00:30 to 23:30 UTC. Total
177 IAT calculated based on the approaches defined by Xu et al. (2020) and Chakraborty et al. (2021)
178 using the following formula.

$$179 \text{IAT}_n = \sqrt{(\text{IATU})_n^2 + (\text{IATV})_n^2}$$

180 Here, n denotes each of the aerosol species treated, whereas U and V indicate corresponding zonal
181 and meridional components of vertically integrated aerosol mass fluxes. We selected U and V
182 components of each species namely BC, OC, and dust.

183

184 2.3. HYSPLIT model

185 To investigate the transport pathways and further provide evidence of the source-sink
186 footprint of different pollutants. We calculated the seven-day back-trajectories (forward and
187 backward in time) using the HYSPLIT model. Trajectories were initialized at 100 m above ground
188 level for the 6 h interval (UTC, 00:00, 06:00, 12:00, and 18:00) at Langtang (28.21°N , 85.61°E ;
189 4900 m a.s.l) in the southern slope of the Himalaya considering it as the transition point between
190 polluted South Asia and background region (i.e., Tibetan Plateau or the Himalaya). The model was
191 run for the entire year using gridded meteorological reanalysis dataset Global Data Assimilation
192 System (GDAS, 1° spatial resolution) (<https://www.ready.noaa.gov/gdas1.php>).

193 2.4. Emission, observation, and reanalysis datasets

194 The anthropogenic emissions are obtained from Emission Database for Global
195 Atmospheric Research and Hemisphere Transport of Air Pollution (EDGAR-HTAP, v2.2) which
196 has been compiled using regional emissions grid maps for BC, OC, $\text{PM}_{2.5}$, PM_{10} , CH_4 , SO_2 , CO,
197 NO_x , and NH_3 from different emissions sectors like industry, residential, agriculture, power,
198 transportation, aviation, and shipping at $0.1^\circ \times 0.1^\circ$ (Janssens-Maenhout et al., 2015). Similarly, the



199 biogenic emissions are prepared from the Model of Emission of Gases and Aerosol from Nature
200 (MEGAN) (Guenther et al., 2006). Additionally, biomass burning emissions are obtained from the
201 Fire Inventory NCAR (FINN) from the National Center for Atmospheric Research (NCAR) with
202 hourly temporal resolution and 1 km horizontal resolution (Wiedinmyer et al., 2011).

203 To examine the model simulated vs. surface concentrations of $PM_{2.5}$ and PM_{10} , we obtained
204 the observational data from the China Environmental Monitoring Stations accessed via
205 (<https://quotsoft.net/air/#archive>). The daily average data from eight-station in mainland China
206 were used to validate against model simulation.

207 For comparison, we also obtained the Aerosol Robotic Network (AERONET)
208 (<https://aeronet.gsfc.nasa.gov/>) Level 2 AOD (aerosol optical depth) data at the locations namely
209 Beijing (40.00°E, 116.38°N) in China, QOMS (28.37°E, 86.95°N, and Nam Co (30.77°E, 90.96°N)
210 in Tibet (China), Langtang (28.21°E, 85.61°N) and Lumbini (27.49°E, 83.28°N) in Nepal, Dushanbe
211 (38.55°E, 68.86°N) in Tajikistan, Mezaira (23.15°E, 53.78°N) in UAE, Dhaka (23.73°E, 90.40°N) in
212 Bangladesh, Kanpur (26.51°E, 80.23°N) in India, Karachi (24.87°E, 67.03°N) in Pakistan. Karachi,
213 Kanpur, Lumbini, and Dhaka are located in IGP. Level 2.0 data comes with cloud-screened and
214 assured quality (with uncertainty range 0.01-0.02) but lack in consistency opted further treatment,
215 which we have constrained by Level 1 data in our study. More details about calibration, data
216 processing, instrumentation, accuracy assessment, and uncertainty are mentioned elsewhere (Eck
217 et al., 1999; Holben et al., 1998; Dubovik et al., 2000).

218 The Copernicus Atmospheric Monitoring Service (CAMS) AOD was obtained from
219 (<https://ads.atmosphere.copernicus.eu/>) which is the fourth generation Atmospheric Composition
220 Reanalysis (EAC4) (Inness et al., 2019). The Visible Infrared Imaging Radiometer Suite (VIIRS)
221 AOD was obtained via (<https://ladsweb.modaps.eosdis.nasa.gov/>). Additionally, MERRA-2 AOD
222 and dust were obtained from Goddard Earth Sciences Data and Information Services Center (GES
223 DISC, <https://disc.gsfc.nasa.gov/>) which have 0.25×0.25 spatial resolution.

224 **2. 3. Results and discussion**

225 **3.1. Comparative assessment of $PM_{2.5}$, PM_{10} , and AOD**

226 To provide an insight into the model reproducibility here we strengthen our study by
227 performing a comprehensive assessment by comparing model outputs with observations for
228 selected pollutants. In doing so, we have shown PM_{10} and $PM_{2.5}$ concentrations in 2017 from
229 measurements and corresponding WRF-Chem simulation at eight locations over mainland China

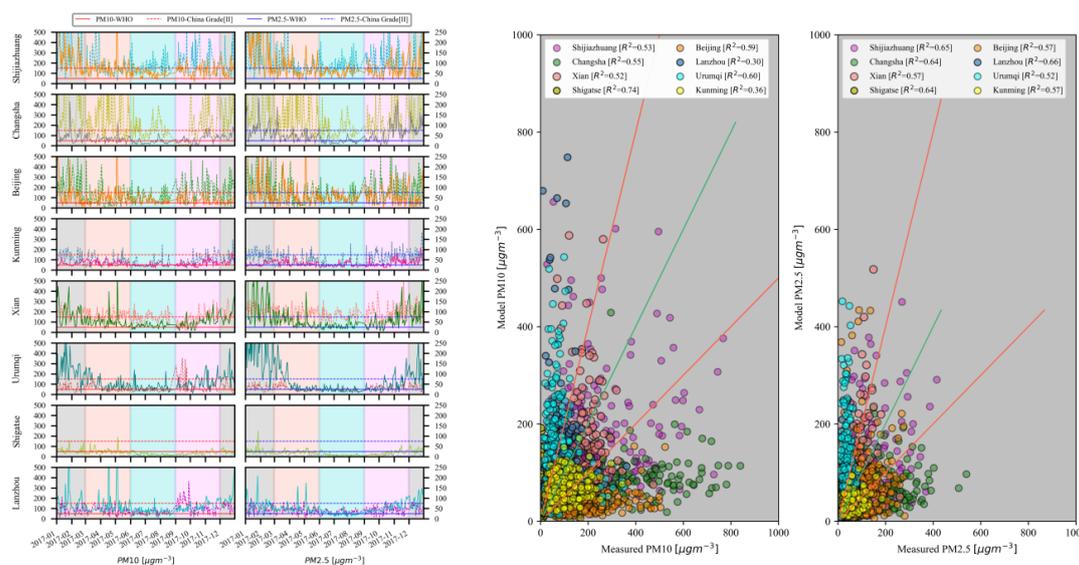


230 in Figure 2. Amongst the selected sites, measured annual mean PM_{10} shows low values in
231 background sites [Shigatse ($41 \mu\text{g m}^{-3}$)] whereas higher concentrations were recorded in urban
232 locations [e.g., Shijiazhuang ($149 \mu\text{g m}^{-3}$), Xi'an ($112.4 \mu\text{g m}^{-3}$), and Beijing ($98 \mu\text{g m}^{-3}$)]. The
233 $PM_{2.5}$ concentrations hold the same patterns as in the case of PM_{10} , relatively low concentrations
234 were observed in a pristine location like Shigatse ($18.3 \mu\text{g m}^{-3}$) and higher in Shijiazhuang (86.1
235 $\mu\text{g m}^{-3}$), Xi'an ($66.5 \mu\text{g m}^{-3}$) and Beijing ($62.6 \mu\text{g m}^{-3}$). Model overestimated $PM_{2.5}$ with a mean
236 bias (MB) in Shijiazhuang ($51.9 \mu\text{g m}^{-3}$), Changsha ($170.1 \mu\text{g m}^{-3}$), Beijing ($49 \mu\text{g m}^{-3}$), and
237 Kunming ($22.9 \mu\text{g m}^{-3}$) whereas underestimated (with MB) in Xi'an ($-93.6 \mu\text{g m}^{-3}$), Urumqi (-61.1
238 $\mu\text{g m}^{-3}$), Shigatse ($-34.2 \mu\text{g m}^{-3}$) and Lanzhou ($-36.3 \mu\text{g m}^{-3}$). The frequency of episodic spikes
239 densely occurs during winter and before the onset of monsoon (Figure 2), however, the model is
240 unable to reproduce these peaks. In general, the model captured the temporal variations of
241 particulate matter concentrations fairly well. The R^2 values indicated that the model reproduced
242 the PM_{10} and $PM_{2.5}$ at an acceptable level for all stations except in Lanzhou ($R^2 = 0.40$) and
243 Kunming ($R^2 = 0.36$) for PM_{10} . The bias could be jointly affected by unaccounted factors, such as
244 meteorological fluctuations that the model was unable to capture (Zhang et al., 2016; Ding et al.,
245 2019), anthropogenic and natural emissions (Ukhov et al., 2020b; Cai et al., 2017), and geography
246 (Wang et al., 2018a; Zhao et al., 2020b).

247 In response to the strict government clean air action plan, a significant reduction of PM_{10}
248 and $PM_{2.5}$ over China has been reported (Ding et al., 2019; Zhao et al., 2020b; Uno et al., 2020;
249 Kang et al., 2019) however PM_{10} and $PM_{2.5}$ values are much higher than the guideline values of
250 the World Health Organization (WHO) and national ambient air quality standard, China. As shown
251 in Figure 2, the daily mean PM concentrations at all locations never drop below the WHO limit of
252 PM_{10} ($50 \mu\text{g m}^{-3}$) and $PM_{2.5}$ ($25 \mu\text{g m}^{-3}$) except for a few occasions. During episodic events, the
253 WHO and China limits for PM_{10} and $PM_{2.5}$ were exceeded by a factor of 10 and 3 respectively,
254 which is a matter of concern to be considered while formulating air pollution abatement policies
255 and actions plans.

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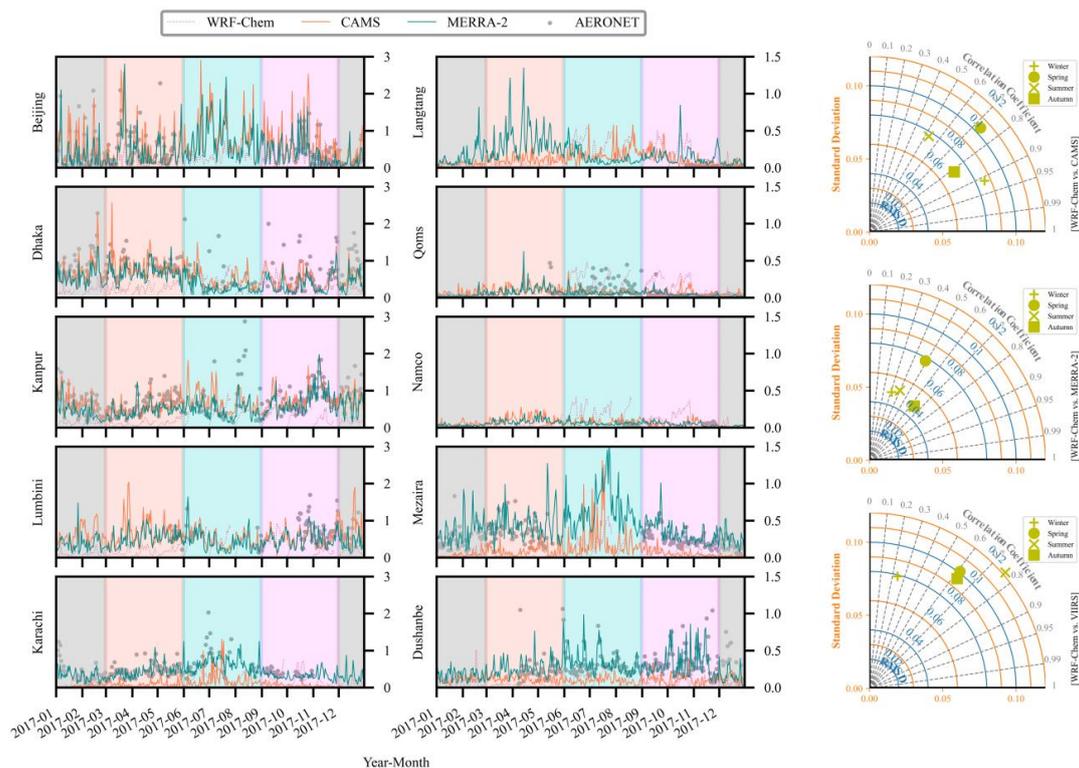


258
 259 Figure 2. Daily PM₁₀ and PM_{2.5} surface mass concentration for respective given locations from
 260 the observations and the corresponding WRF-Chem simulations for the simulation period. For
 261 each panel, the dashed and solid line corresponds to model and observation data whereas the
 262 horizontal dashed and the solid line indicates the air quality guideline values of WHO and China.
 263 From left to right, different color in the background represents the winter, spring summer, and
 264 autumn seasons. Scatter plot showing model vs. observation PM₁₀ and PM_{2.5}. Different dots
 265 colors represent the location-wise data. Solid straight lines are (2:1) - upper orange (1:1) - green
 266 and (2:1) - lower orange lines. The correlation coefficient values for each location are shown in
 267 the above figure. Other statistical metrics are given in Table S1.

268 AOD, an aerosol proxy, from the background and polluted stations was utilized for model
 269 validation. Figure 3 (first and second) shows a collocated time series plot for the observation,
 270 reanalysis, and modeled AOD whereas the Taylor diagram in the third column illustrated the
 271 seasonal mean statistical metrics for the model versus CAMS, MERRA-2, and VIIRS respectively.
 272 The AOD values correspond to urban than background locations. For all considered AERONET
 273 stations, the statistical score provided in Table 2s illustrated that the ground-based AOD
 274 satisfactory correlated with CAMS and MERRA-2 than WRF-Chem output. It is worth mentioning
 275 here that Gueymard and Yang (2020) study shows that MERRA-2 performs better than CAMS
 276 over most continents climates. The model vs. measured R^2 value was observed to be very low in
 277 Karachi ($R^2 = 0.257$, $RMSE = 0.470$) and Kanpur ($R^2 = 0.278$, $RMSE = 0.470$).



278 The observed mean AOD values in 2017 at AERONET sites in Beijing (0.44 ± 0.13),
 279 Dhaka (0.71 ± 0.14), Kanpur (0.69 ± 0.27), Mezaira (0.27 ± 0.08), QOMS (0.06 ± 0.05), Nam Co
 280 (0.06 ± 0.02), Lumbini (0.57 ± 0.20), Karachi (0.48 ± 0.16), and Dushanbe (0.30 ± 0.05) are
 281 comparable with other studies by Ramachandran et al. (2020a) from IGP locations, Pokharel et al.
 282 (2019) from Nam Co and Qoms, Rupakheti et al. (2018) from Lumbini, Rai et al. (2019) from
 283 Langtang, and Rupakheti et al. (2020) from Dushanbe. In the present study, the simulated AOD
 284 values are in line with previous findings. However, intrinsic uncertainty could be associated with
 285 the model resolution, emission data, meteorological influence, and data availability during the
 286 study period.



287
 288 Figure 3. Time series of daily average AOD for the year 2017 (January-December) at ten
 289 AERONET sites for WRF-Chem (dashed magenta), CAMS (solid yellow), MERRA-2 (solid tan),
 290 and AERONET (circle grey). The first and second column shows the AERONET AOD with
 291 reanalysis data and WRF-Chem results for urban and background locations respectively. In the

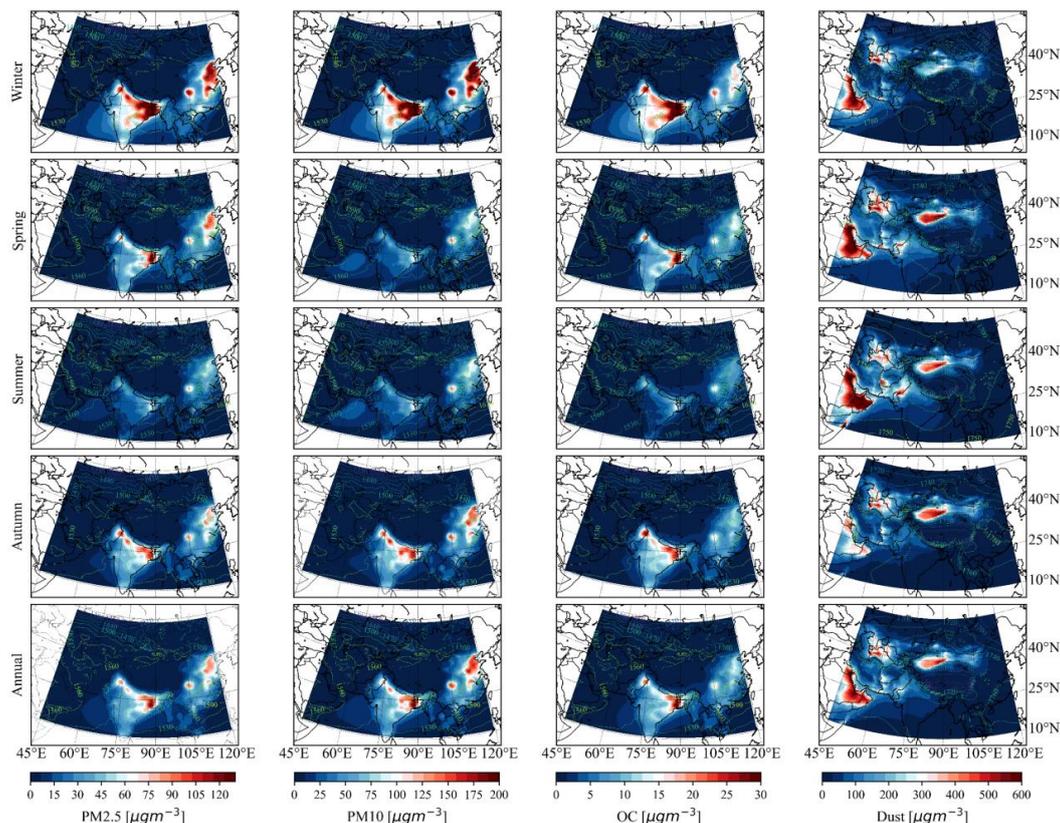


292 last column, the Taylor diagram represents the seasonal statistical metrics of WRF-Chem
293 simulation data against VIIRS, MERRA-2, and CAMS data.

294 3. 3.2. Spatio-temporal seasonal aerosols and AOD variation and meteorological influence

295 The changes in atmospheric dynamics and meteorological conditions besides the emissions
296 can change aerosol concentrations over the PTP region highlighting the need for studies on
297 meteorological influence on spatiotemporal distributions of aerosols. Hence, exploiting model
298 aerosol diagnostic output together with meteorological parameters here we unraveled the effects
299 of such meteorological conditions on the spatiotemporal distribution of aerosol. For $PM_{2.5}$, PM_{10} ,
300 and OC three major hotspots with intense anthropogenic emissions were identified namely IGP,
301 Sichuan Basin, and East China. In contrast, heavy dust concentrations are laden over Saudi Arabia,
302 Central Asia, Taklamakan Desert, Gobi Desert, and Thar Desert (Figure 3, last column). Figure 4
303 shows clear seasonal variability with winter high and summer low concentrations for all aerosol
304 species. In winter, aerosol concentrations were estimated maximum in the range $75\text{--}125\ \mu\text{g m}^{-3}$
305 ($PM_{2.5}$), $100\text{--}200\ \mu\text{g m}^{-3}$ (PM_{10}), $15\text{--}30\ \mu\text{g m}^{-3}$ (OC) over South Asia and East China. During
306 spring, aerosols concentrations were found to be relatively lower than winter ($PM_{2.5}$: $60\text{--}100\ \mu\text{g m}^{-3}$,
307 PM_{10} : $25\text{--}125\ \mu\text{g m}^{-3}$, OC: $10\text{--}25\ \mu\text{g m}^{-3}$) reaching to lowest levels during summer ($PM_{2.5}$: 15--
308 $70\ \mu\text{g m}^{-3}$, PM_{10} : $25\text{--}110\ \mu\text{g m}^{-3}$, OC: $5\text{--}15\ \mu\text{g m}^{-3}$) but again rebound to higher concentrations
309 during autumn ($PM_{2.5}$: $30\text{--}90\ \mu\text{g m}^{-3}$, PM_{10} : $20\text{--}150\ \mu\text{g m}^{-3}$, OC: $15\text{--}20\ \mu\text{g m}^{-3}$).

310 For the $PM_{2.5}$, the results from our simulation are comparable with Ojha et al. (2020) over
311 the IGP region and Zhang et al. (2019a) over Eastern China. The maximum dust concentrations
312 were over the Arabian Peninsula with concentrations of $400\text{--}600\ \mu\text{g m}^{-3}$, followed by the
313 Taklamakan and Gobi Desert ($300\text{--}500\ \mu\text{g m}^{-3}$), Central Asia ($200\text{--}400\ \mu\text{g m}^{-3}$), and Thar Desert
314 ($200\text{--}450\ \mu\text{g m}^{-3}$). Annually, higher concentrations of $PM_{2.5}$, PM_{10} , and OC were estimated over
315 the SA than EC region. Such changing pattern of aerosols concentrations over the region is also
316 reported in previous studies (Crippa et al., 2018; Ding et al., 2019; Uno et al., 2020; Kanaya et al.,
317 2020; Zhao et al., 2020c). In contrary to $PM_{2.5}$ and PM_{10} , OC concentrations were estimated higher
318 over SA than EC as the emissions from biomass burning contributed more OC over SA than
319 Eastern China. As stated by Han et al. (2020), extremely strong biomass burning in SA, mostly in
320 spring affects the BC and OC concentrations.



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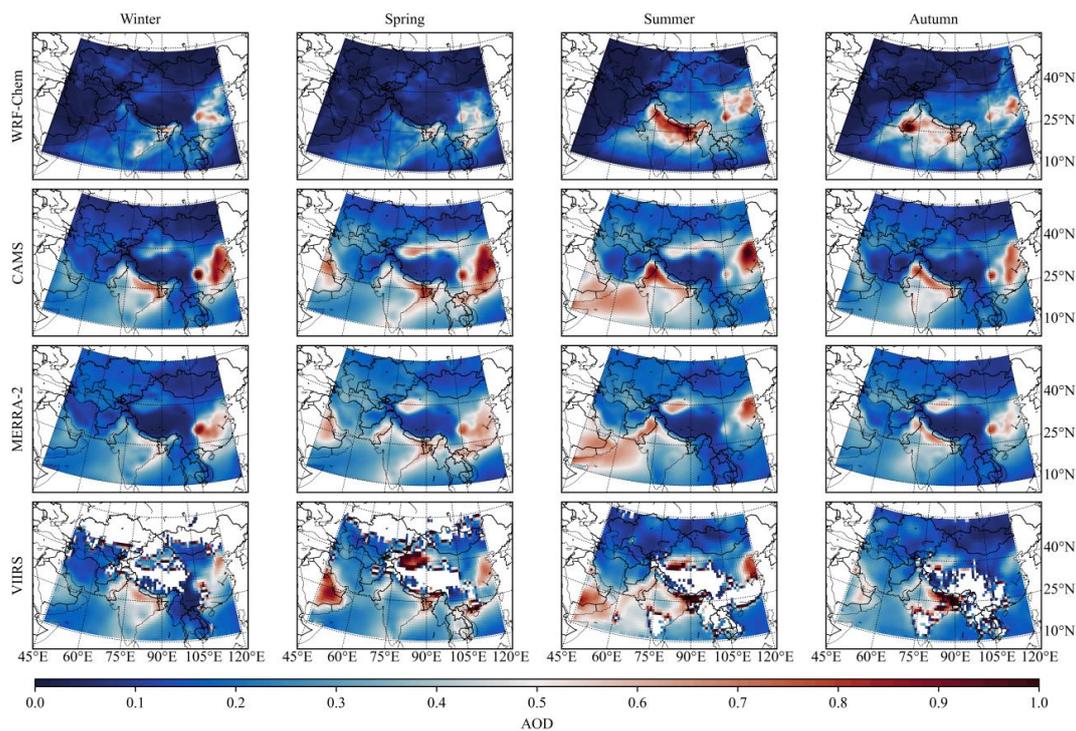
322 Figure 4. Seasonal spatio-temporal variation of aerosols ($PM_{2.5}$, PM_{10} , OC, and Dust). Color-coded
323 contour line represents geopotential height at 850 hPa.

324 Regarding the AOD, as shown in Figure 5, a large spatial variation in AOD is observed
325 over the PTP region. The higher AOD (>0.6) dominates over aerosols pool regions (Figure 4), i.e.,
326 IGP region, East India, Bangladesh, Sichuan Basin, East China, and dust rich regions like Saudi
327 Arabia, Central Asia, Taklamakan Desert, Gobi Desert, and the Thar Desert. These regions are
328 known to have heavy atmospheric aerosol loadings from anthropogenic emissions, wildfire
329 emissions, biomass burning, and dust emissions (Gautam et al., 2011; Chen et al., 2013). On
330 contrary but as expected, a relatively pristine region (i.e., TP, and the Himalayas) has lower AOD
331 values as showed by WRF-Chem model, CAMS, MERRA-2, and VIIRS data throughout the
332 simulation period. Our simulation results showed that the AOD was underestimated compared to
333 the reanalysis data and satellite products. However, zonal mean statistical metrics (Figure 3, Taylor



334 diagram) indicated that simulated results captured the general spatial-temporal patterns to some
335 extent. Indeed, it should be noted that WRF-Chem masked out the general feature of AOD over
336 Taklamakan Desert and Gobi Desert their surrounding regions. Over the Middle East and Arabian
337 Peninsula, the AOD was underestimated by a factor of two. Such bias could be due to the choice
338 of parameterization and the function used in the GOCART scheme where dust-dominated source
339 regions are not well reproduced as the model is forced to run with a high spatial resolution (30
340 km). Uncertainty in AOD representation is in GOCART model described elsewhere (Parajuli et
341 al., 2019; Ukhov et al., 2020a; Zhao et al., 2013). From a regional perspective, it is evident that
342 the high magnitude of AOD is closely associated with the intense emission source regions
343 including IGP, East China, and Sichuan Basin in all seasons that could have been resulted from
344 rapid economic development including urbanization and industrial expansion. Another intrigued
345 factor could be associated with biomass burning and anthropogenic emissions from highly
346 populated regions. The distinct feature of high AOD over the Sichuan Basin could collectively be
347 attributed to emission sources, topography, and meteorological conditions. As the mountain ridges
348 inhibit the dispersion and transport of pollutants and stagnant meteorological conditions restrict
349 the advection process that ultimately causes aerosol accumulation over a bowl-shaped basin which
350 could result in higher AOD (Fan et al., 2020; Cao et al., 2020). Interestingly, high AOD also
351 occurred over the Persian Gulf and the Arabian Sea. It could be linked with dust outflow from dust
352 source regions in the Middle East in the West and the Thar Desert in the East (Figure 4). As stated
353 by Nabavi et al. (2017) and Parajuli et al. (2019) the effect of the prevailing Shamal winds (strong
354 northwesterly winds in the region) could result in dust storms, increased wind-blown dust, poor
355 visibility, and worse air quality, and elevated AODs over the surrounding regions.

356 Irrespective of the seasonality, summer high AOD was found to be along the IGP region.
357 A similar pattern with high AOD in summer and low in winter over IGP was reported in previous
358 studies as well (Kulkarni et al., 2015; Ukhov et al., 2020a; Wang et al., 2020). In addition, a
359 comprehensive study on multi-model evaluation by Pan et al. (2015) showed a clear indication of
360 summer high and winter low over South Asia, with the maximum AOD stretching westward from
361 East India to Kanpur in central IGP to Lahore in the western IGP and thus covering the entire IGP
362 region.



363

364 Figure 5. The seasonal mean AODs from the model simulation, reanalysis data, and satellite
365 product for the simulation period (2017).

366 According to Ratnam et al. (2021) such enhanced AOD during summer is partly due to
367 aerosol long-range transport through the low-level jet and tropical easterly jet that persists over the
368 IGP region and due to a greater number of break spells. As stated by Pan et al. (2015), such AOD
369 underestimation during winter over the IGP region might have been due to underlying multiple
370 factors such as wintertime relative humidity (RH), proportionally higher presence of nitrate
371 component in aerosols, inadequate representation of anthropogenic and biofuel emission, and open
372 biomass burning activities. Feng et al. (2016) stated that the humidity bias results in the
373 underestimation of AOD prediction over the South Asia region. Thus, summer high AOD could
374 be modulated by enhanced temperature and RH that intensify hygroscopic growth of aerosols,
375 consequently yielding high AOD (Alam et al., 2012; Altaratz et al., 2013). The biomass burning
376 emissions, notably from forest fires and agro residue burning especially during the winter and
377 spring seasons are a dominant source of pollutants and hence enhanced AODs over South Asia



378 (Ramachandran et al., 2020b). Thus, it is believed that specific biomass burning events could not
379 be captured well by the model subsequently led to an underestimation of AOD.

380 Emissions under unfavorable meteorological conditions for their dispersion cause air
381 pollution episodes. Thus, apart from emissions, the variabilities in concentrations of air pollutants
382 must depend on the effects of meteorological conditions which are of great significance in
383 determining dispersion, atmospheric processing, and removal of air pollutants (Wang et al.,
384 2018b). To showcase meteorological influence on the variation of aerosols here we have presented
385 seasonal and annual meteorological parameters (i.e., total precipitation, planetary boundary layer
386 height (PBLH), RH, and wind speed (Figure 6). The minimum aerosol concentrations were
387 observed over IGP, East China, Bay of Bengal, and Arabian Peninsula (Figure 4) during summer
388 coinciding with the maximum precipitation (Figure 6). The onset of the summer monsoon (rainy
389 season) facilitates the wet removal process. Additionally, dynamical change in PBLH across South
390 Asia and East China during summer further helps reduce aerosol concentrations near the surface
391 significantly due to more efficient dispersion and transport of aerosols. A study by Zhao et al.
392 (2020a) stated that the elevated PBLH facilitates vertical diffusion, leading to a reduction in
393 PM_{2.5} concentration over the Sichuan Basin. On contrary, aerosol concentrations typically increase
394 during winter as it is favored by less precipitation, shallow PBLH (<100 m), low wind speed over
395 extended periods, and additional seasonal sources such as brick production in South Asia (Mues
396 et al, 2018). The seasonal simulated wind speed is relatively lower over the highly polluted region
397 (i.e., South Asia) that inhibited the dispersion and transportation of aerosols. However, the high
398 wind speed found over Arabian Peninsula that could drive dust storms and transport dust from
399 such desert environments to downwind regions. Zhang et al. (2019b) study from East China and
400 Ojha et al. (2020) study from South Asia also stated that haze events were associated with severe
401 air stagnation conditions represented by low wind speed, low PBLH, temperature reduction, and
402 high humidity.

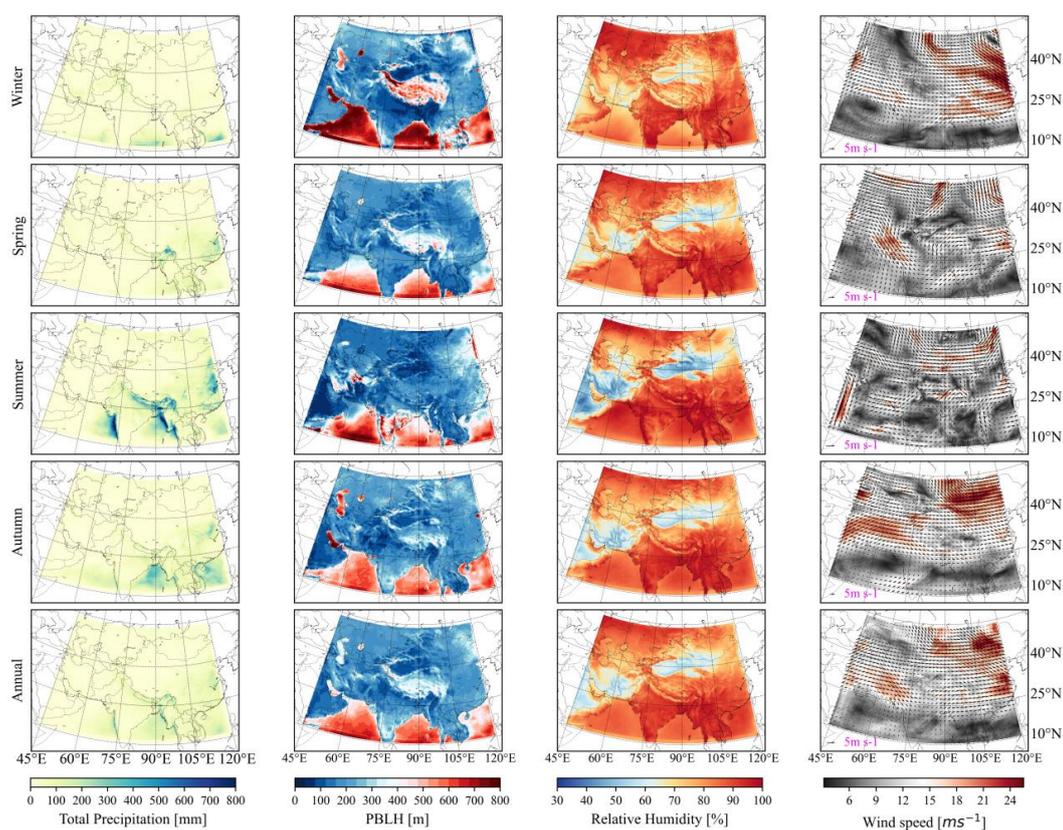
403

404 **3.3. Aerosols transport dynamics**

405 The interaction between meteorological conditions and terrain exacerbates the severity of
406 air pollution (Wang et al., 2018b). Under favorable circumstances, the aerosol particles are
407 dispersed well and transported afar from source regions more efficiently. Thus, to examine such a
408 mechanism and offer insight into the vertical distribution of aerosols here we have provided the



409 cross-sectional zonal mean concentrations (Figure 7). The seasonal columnar profile exhibited that
410 the maximum aerosols concentrations are limited within 1-2 km over South Asia and East China
411 regions. The highest PM_{2.5}, PM₁₀, and OC concentrations are found near the surface with
412 concentrations in the range of 50-100 $\mu\text{g m}^{-3}$, 70-100 $\mu\text{g m}^{-3}$, and 10-20 $\mu\text{g m}^{-3}$, respectively over
413 the high emission source regions like IGP in the southern flank of the Himalayas and East China.
414 Whereas, the higher extinction coefficient values correspond with the aerosol concentrations.

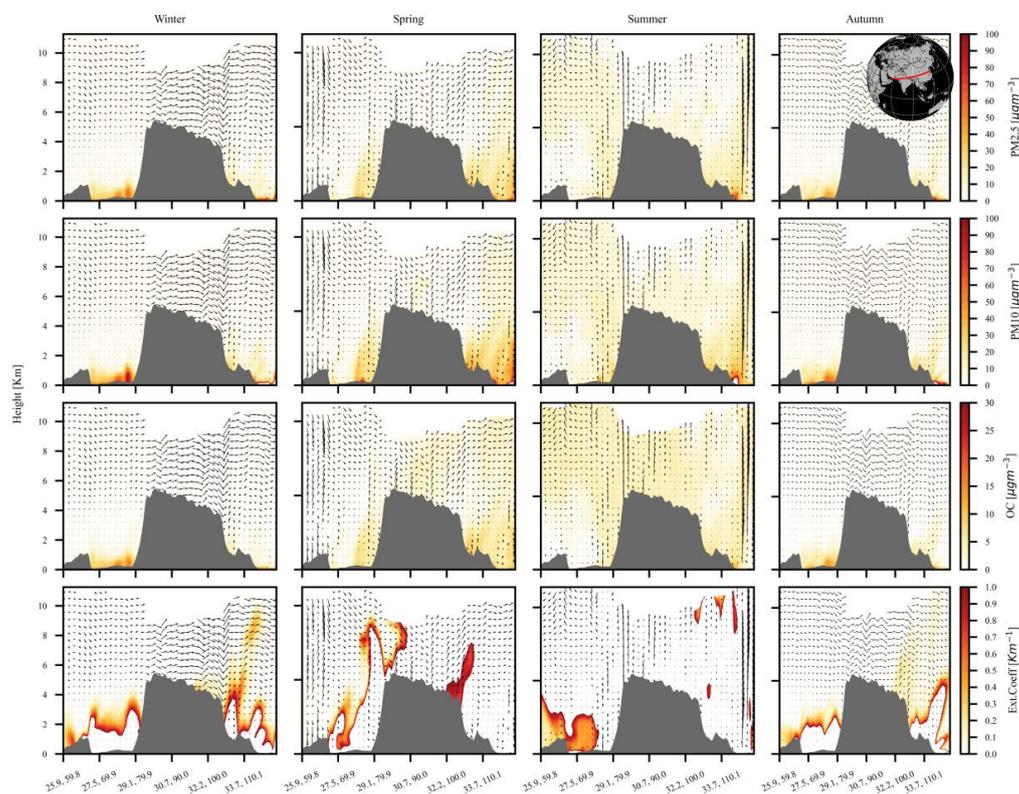


415
416 Figure 6. WRF-simulated seasonal and annual mean meteorological parameters. Wind speed
417 interpolated at 500 hPa is indicated by color-code.

418 The maximum values of extinction coefficient were observed over aerosol-rich regions,
419 i.e., South Asia and East China. However, noticeable higher values during winter and spring
420 seasons over the HTP can be attributed to higher amounts of aerosols transported across the
421 Himalayas to the Tibetan Plateau from upwind source regions in South Asia (Lüthi et al., 2015).



422 During winter, higher aerosol concentrations are confined near the surface. However, aerosols start
423 getting into higher elevations during spring. Such evident accumulation of aerosols was also
424 observed by Zhang et al. (2020a). Apart from emissions, this could be attributed to weakening
425 winter monsoon towards the end of the winter season, high humidity, an increase in the occurrence
426 of calmer winds, reduction of wind shear of horizontal zonal winds, a decrease in precipitation,
427 and low PBLH (Chen and Wang, 2015; Zhang et al., 2016).



428
429 Figure 7. Seasonal cross-sectional columnar profile for PM_{2.5}, PM₁₀, OC, and aerosol extinction
430 coefficient. The red line in the upper right corner shows the starting and endpoints of the cross-
431 section transect. The filled gray color indicates terrain height and the arrow represents wind speed
432 and wind direction.

433 In contrast, during summer the aerosol concentrations are elevated vertically as high as up
434 to 10 km potentially driven by the deep convection and other summer monsoon activities. During
435 summer in Figure 8, the higher density of air masses at 800-100 hPa in the Himalayas are received



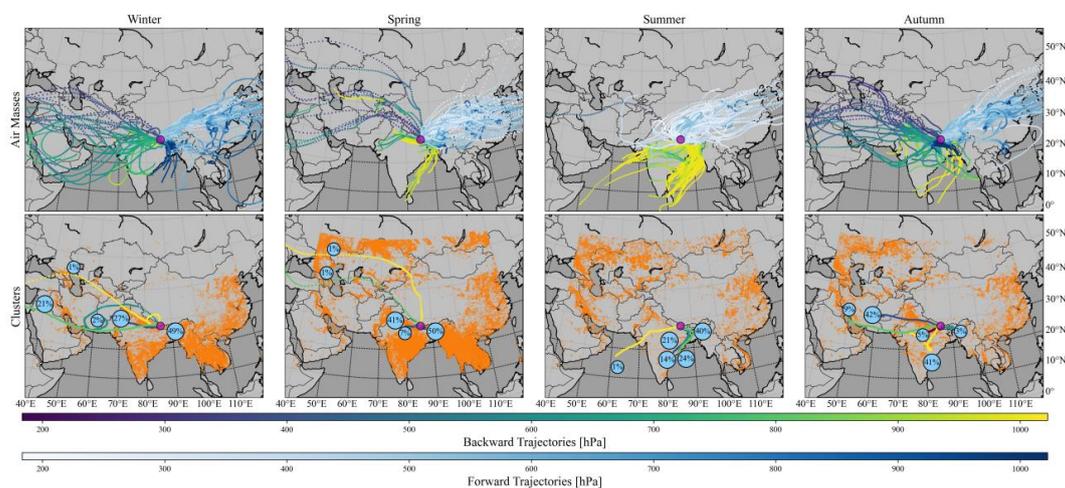
436 from the Bay of Bengal and Arabian Peninsula which further moved forward beyond the
437 Himalayas and TP at 200-400 hPa. Such elevation differentiates air masses that could carry aerosol
438 species from the surface to the different atmospheric layers. Bian et al. (2020) stated that
439 convective overshooting events are also an efficient mechanism to transport the aerosols to the
440 upper troposphere and lower stratosphere (UTLS) within the Asian Summer monsoon anticyclone.
441 According to Zhao et al. (2017), the BC transport from South Asia to the TP during the summer
442 monsoon was found to be strongly dependent on cyclonic activities and convergent airflows which
443 significantly increased the BC in the TP. Whereas, Rai et al. (2022) stated that the role of different
444 wind regime plays a crucial for the BC transportation that is even injected into UTLS region. As
445 autumn approaches, the summer monsoon withdrawal process continues thereby weakening the
446 convection process, which ultimately reduced the vertical aerosols concentrations.

447 Further, to investigate the aerosol transport mechanisms and pathways here we traced back
448 air mass trajectories and assessed their forward trajectories, including through a clusters analysis
449 (Figure 8). Except for the summer season, westerlies are a predominant wind received in the
450 Himalayas (at Langtang) which moved forward (mostly eastward) to the Tibetan Plateau and East
451 China. The clusters analysis indicated that the higher percentage of air masses traversed long
452 distances over land (22% as far back as the Middle East and Central Asia and 30% confined
453 mostly within South Asia) in long and dry winter and spring seasons that potentially transported
454 higher aerosol concentrations from upwind intense source regions through long-range transport
455 mechanism. On the other hand, the Bay of Bengal is the main source of air masses (mostly carrying
456 heavy amounts of moisture) during summer that could lower aerosol concentrations at the surface
457 through the wet scavenging process. However, such air masses could also inject pollution into the
458 higher layers through deep convection and orographic lifting processes over the Himalayan region
459 (Pan et al., 2015; Bian et al., 2020; Zhang et al., 2015b).

460 Previously, numerous studies have explored and documented the characteristics and
461 mechanisms of aerosol transport over the HTP region (Han et al., 2020; Kopacz et al., 2011; Lüthi
462 et al., 2015; Yang et al., 2018; Zhang et al., 2015a). Despite these findings, the current
463 understanding of aerosol transport mechanisms and pathways, particularly its seasonal extremes is
464 yet to be explored and advanced. Over the period, there has also been significant growth in studies
465 related to water vapor transport that was termed as an atmospheric river (AR) (Neiman et al., 2011;
466 Ralph et al., 2017; Guan and Waliser, 2015). Further, Chakraborty et al. (2021) extended the AR



467 concept in terms of aerosol atmosphere river (AAR). Voss et al. (2020) investigated the dynamic
468 association between dust and AR. The integrated aerosol transport (IAT) studies over the PTP
469 region are rather recent. Therefore, here we leverage recent advances in IAT detection algorithm
470 (Xu et al., 2020) to MERRA-2 dataset (Gelaro et al., 2017) to investigate spatiotemporal IAT
471 distribution and potential association with the meteorology of the region.

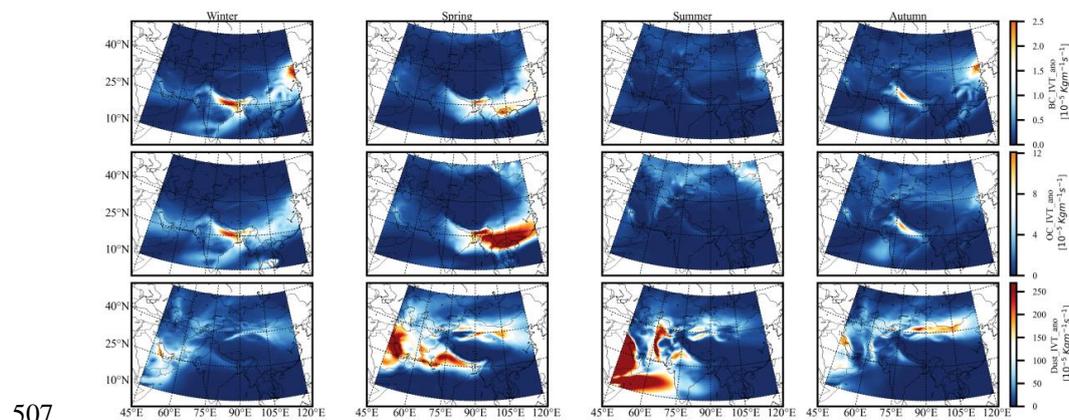


472
473 Figure 8. Seven-day seasonal backward/forward air mass trajectories and air mass clusters at
474 Langtang (Nepal) for the year 2017 were generated by the HYSPLIT model. The yellow dotted
475 color in the background represents the distribution of fire spot events that occurred in 2017. The
476 fire data was obtained from the Suomi National Polar-Orbiting Partnership (Suomi NPP) satellite
477 platform, which was acquired by the Visible Infrared Imaging Radiometer Suite (VIIRS).

478 During winter, BC_IAT showed a bipolar spatial distribution having one maximum over
479 IGP in South Asia and another maximum over North China Plain (NCP) in East Asia, with the
480 maximum BC_IAT of $2\text{--}2.5 \times 10^{-5} \text{ kg m}^{-1} \text{ s}^{-1}$. During the spring season, the BC_IAT pattern shifts
481 eastwards, particularly the South Asian BC-IAP extending to northern Southeast Asia and south
482 East Asia. The weakest BC_IAT was observed over both IGP ($\sim 1 \times 10^{-5} \text{ kg m}^{-1} \text{ s}^{-1}$) and EC ($1.5 \times$
483 $10^{-5} \text{ kg m}^{-1} \text{ s}^{-1}$) during summer. Nevertheless, during autumn BC_IAT rebounded to its maximum
484 of about $1\text{--}2 \times 10^{-5} \text{ kg m}^{-1} \text{ s}^{-1}$ over IGP and $1.5\text{--}2 \times 10^{-5} \text{ kg m}^{-1} \text{ s}^{-1}$ over EC. A similar spatiotemporal
485 pattern has been observed for the OC_IAT. During winter, the IGP region had higher OC_IAT (8--
486 $10 \times 10^{-5} \text{ kg m}^{-1} \text{ s}^{-1}$) than the rest of the region. In spring, the OC_IAT pattern shifted toward East
487 Asia along $90^\circ\text{E}\text{--}120^\circ\text{E}$ covering a large swath of northern Southeast Asia with very high OC_IAT
488 (Figure 9). A clear picture of OC_IAT spring anomaly could be seen over East Asia (Figure 10),



489 with the extreme reaching a maximum ~ 5 folds higher than the annual average (Figure 10). This
490 could be explained by the contribution from the biomass burning emission. During spring, the fire
491 events are maximum over Asia (Figure 8, spring). It can be thus inferred that OC_IAT contributed
492 significantly to increasing aerosol extinction coefficient even over the HTP region (Figure 7).
493 Chakraborty et al. (2021) mentioned that higher population regions like China and India are
494 associated with BC and OC AAR that transport BC and OC aerosols either in the direction of the
495 trade winds in tropics or westerlies in the multitude. Unlike BC_IAT and OC_IAT, a contrasting
496 spatial distribution for Dust_IAT was found with a maximum over arid land and deserts including
497 AP, PG, CA, TD, Gobi Desert, and the Thar Desert. Seasonally, the highest Dust_IAT ($200\text{--}280 \times$
498 $10^{-5} \text{ kg m}^{-1} \text{ s}^{-1}$) is seen during summer over part of Saudi Arabia, CA, PG, and IGP. Chakraborty
499 et al. (2021) found the largest Dust_IAT over the Middle East, Central Asia, and East China
500 regions ($\sim 30 \times 10^{-3} \text{ kg m}^{-1} \text{ s}^{-1}$). Both simulation and other datasets revealed that summer high AOD
501 over IGP (Figure 5). Thus, this could be possibly due to the response of high Dust_IAT from the
502 neighboring regions under favorable meteorological conditions. High dust concentrations were
503 found over Arabian Peninsula with the swatch of high dust concentrations extending over the IGP
504 region during winter. However, on a global scale, Chakraborty et al. (2021) revealed that the AAR
505 is responsible for transporting more than 40 % of the total IAT over tropical and mid-latitude
506 regions.

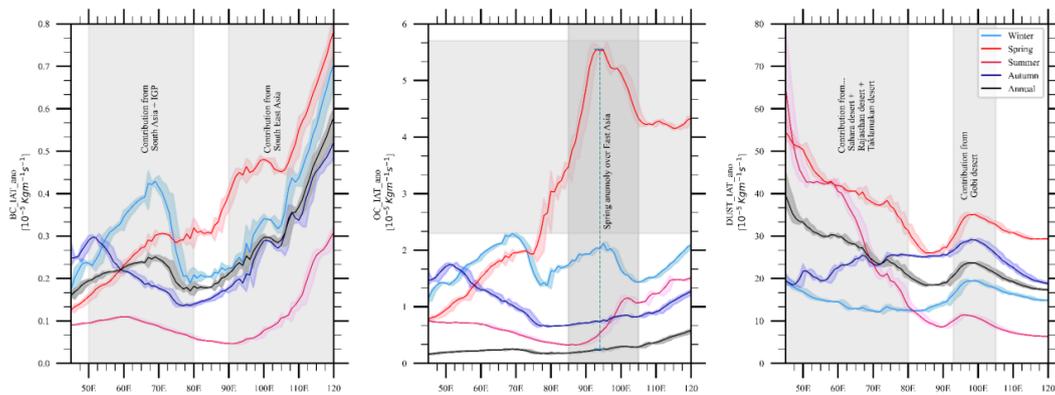


507
508 Figure 9. Seasonal integrated aerosol transport (IAT) for the BC, OC, and Dust for the year 2017
509 over the study region.

510
511



512 Aerosol concentrations and distribution are complex processes governed by emissions,
513 meteorological conditions, and land-atmosphere interactions. Except for emissions, the
514 meteorology of a region is one of the prominent aerosol-modulating factors for aerosol transport
515 and dispersion. Therefore, here we investigate the relationship of IAT with precipitation and wind
516 speed (Figure 11). The strong negative correlation found over the BC and OC emission source
517 regions (i.e., SA and EC) indicated that the precipitation affects the IAT significantly, as higher
518 precipitation is likely to lower the IAT through the wet scavenging of aerosols. On contrary, a
519 positive correlation with precipitation was found over Central Asia and Mongolia where emission
520 is relatively low. Interestingly, the relationship of Dust_IAT with wind speed was opposite to that
521 between IAT and precipitation.

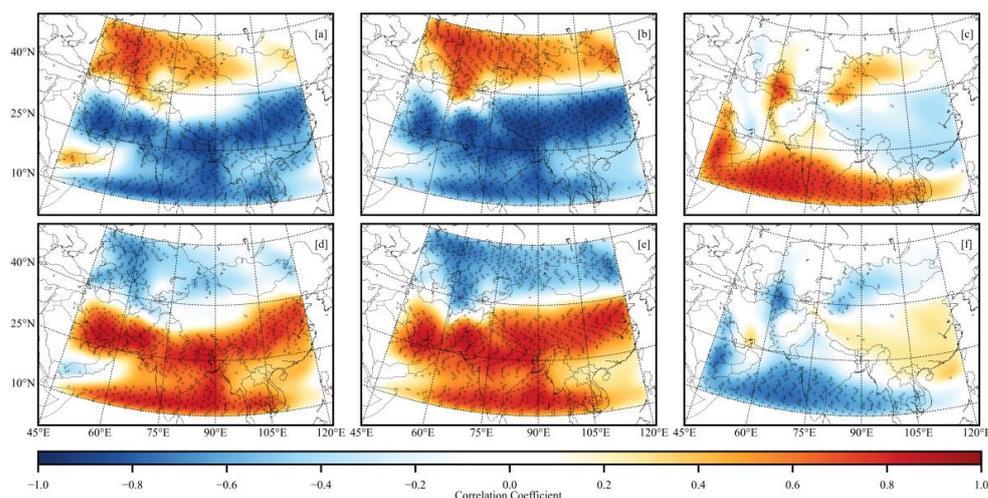


522
523 Figure 10. Seasonal-annual zonal mean of integrated aerosol transport (IAT) of BC, OC, and dust
524 for the year 2017. The solid line indicated the meanwhile shaded identical color represent the $\pm 1\sigma$.

525 Contrary to precipitation, a positive correlation between wind and BC_IAT and OC_IAT
526 were found over the emission source regions (Figure 11, (d, e)). The wind, one of the prominent
527 meteorological drivers plays a crucial role in aerosols transport. The higher the wind speed the
528 more BC and OC carried throughout the atmospheric column. Whereas, Voss et al. (2020) found
529 that the dusty AR occurs when strong winds carry dust from Asia across the Pacific. Thus, it can
530 be inferred that the different meteorological parameters and atmospheric dynamics can offset the
531 benefits availed from a reduction in anthropogenic emissions (Ojha et al., 2020).

532

533



534
535 Figure 11. The relationship of BC_IAT, OC_IAT, and Dust_IAT with precipitation (a, b, c) and
536 wind speed (d, e, f), respectively. Plus (+) symbol represents that passed the significant test with
537 $p < 0.1$.

538

539 Conclusions

540 This study reports the quantification of seasonal aerosol concentrations, investigation in
541 transport dynamics, and meteorological influence over the PTP region by leveraging multi-sensors
542 satellite, reanalysis dataset, and atmospheric model simulations. We have shown that the model
543 can reproduce fairly well $PM_{2.5}$, PM_{10} , and AOD when compared with the observation data.
544 However, we argued that the plausible biases could be due to unfavorable meteorological
545 conditions, emissions, and the geographical setting of the region. Our finding showed that the
546 aerosol concentrations exhibited clear seasonality with winter high and summer low. Such
547 pronounced difference in aerosol concentrations could be favored by less precipitation, shallow
548 PBLH, and low wind speed in winter. The higher AOD values correspond to aerosol
549 concentrations. However, the higher AOD values during summer found over IGP could be
550 explained by enhanced temperature and RH that support the hygroscopic growth of the aerosols.
551 The seasonal columnar profile showed that the maximum aerosol concentrations are limited within
552 2 km above the surface over intense emission source regions like South Asia and East China.



553 However, significant aerosol concentrations and aerosol extinction were also found over the HTP
554 that could be attributed to injection through deep convection, orographic lifting, and long-range
555 transport processes. From IAT calculation, BC and OC transport was confirmed predominantly
556 during winter and spring over emission-rich regions whereas dust transport was originated in arid
557 land and deserts and carried forward (eastward) by westerlies. Interestingly, a clear signature of
558 OC transportation was noticed over Southeast Asia during spring that is directly linked with forest
559 fire events in southern Asia. The spring anomaly of OC_IAT was found to be exceptionally high,
560 five folds higher than the annual average. Additionally, we found that the effect of precipitation
561 on IAT (higher precipitation - lower IAT) is opposite to that of wind speed on IAT (higher wind
562 speed – higher IAT) over the emission source region. It is plausible that low wind speed caused
563 stagnant aerosols but it higher wind speed facilitated transportation and dispersion of aerosols.
564 Thus, we concluded that the distinctive aerosol seasonality and transport dynamics depend on the
565 strength of emission source, process, and governing meteorological drivers. In the future, one
566 could emphasize the critical role of the aerosol atmospheric river in air quality by considering more
567 meteorological variables and societal impact, especially during extreme events.

568

569

570 **Code and data availability**

571 The AOD data used in this study are available at CAMS (<https://ads.atmosphere.copernicus.eu/>),
572 MERRA-2 (<https://disc.gsfc.nasa.gov/>), and VIIRS (<https://ladsweb.modaps.eosdis.nasa.gov/>).
573 The particulate matter data can be obtained via (<https://quotsoft.net/air/#archive>). The
574 observational AOD can be accessed from (<https://aeronet.gsfc.nasa.gov/>). The model used in this
575 study can be accessed from (<https://ruc.noaa.gov/wrf/wrf-chem/>). The other datasets and codes
576 used in this study can be obtained from
577 (https://github.com/mukeshraeee/Mukesh_ACP_Manuscript).

578

579 **Author contributions**

580 MR performed the model simulation, analysis, and prepared the figures. JY, XC, and YH
581 contributed to data curation and formal analysis. SK, MR, DR, LT, and JY supervised writing,
582 review, and editing. SK, JY, and MR conceptualized the paper. MR wrote the paper, with all the
583 contributions from all co-authors. SK supervised the project.



584

585

586 **Competing interests**

587 The contact author has declared that neither they nor their co-authors have any competing interests.

588

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599 and IPART for IAT calculation (Xu et al., 2020).

600

601

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